

# Precipitation from shallow convection, microphysics evaluation, and RACMO turbulence in HARMONIE-AROME-MetCoop

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This poster contains three projects: 1) A parameterization for variable maximum cloud-depth of shallow convection in EDMF-m, 2) Evaluation of cloud liquid and ice against ground based remote sensors. 3) Results from the RACMO turbulence scheme in MetCoop.

## 1. Shallow convection

The study was motivated by lack of precipitation (snow) falling along the northern coast of Norway when shallow convection was present.

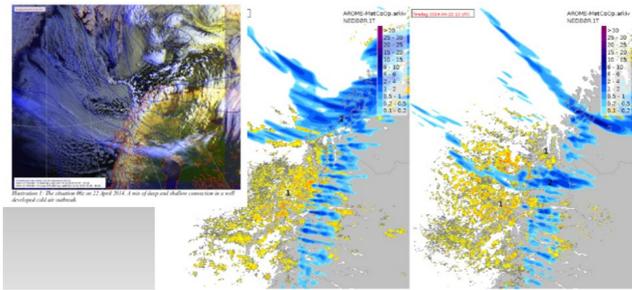


Figure 1: Harmonie-MetCoop 1 h accumulated precipitation compared with Andøya radar (yellow) and satellite (Gunnar Noer, Met).

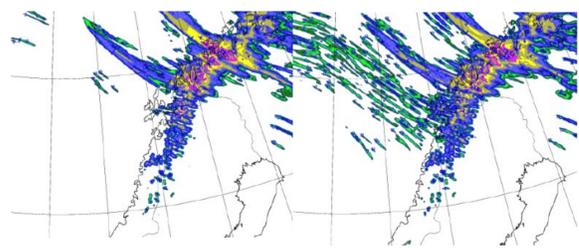


Figure 2: 6 h acc precip. Left, EDMF-m, Right, EDKF

In EDMF-m, if the convective cloud thickness is 4 km or less, it is considered shallow, and only very little precipitation can be generated. (Anything above 4 km is considered deep convection and is not parameterized). However significant snowfall can come from clouds that are only around 1-2 km deep.

Precipitation production is likely to be favoured by active ice-phase processes, so that when cloud-base temperature is close to 0 C, precipitation is possible with relatively shallow convective clouds.

Following Kain, 2004 a variable maximum cloud thickness for shallow convection was introduced:

$$\begin{aligned}
 D &= 4000 & \text{TLCL} > 20 \text{ C} \\
 D &= 1200 & \text{TLCL} < 0 \text{ C} \\
 D &= 1200 + 140\text{TLCL} & 0\text{C} \leq \text{TLCL} \leq 20\text{C}
 \end{aligned}$$

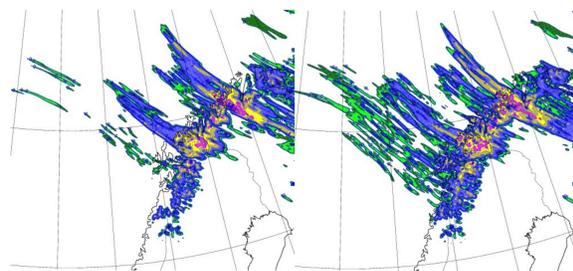


Figure 3: 6 h acc precip. Harmonie-AROME-MetCoop, Left reference 38h1.2, right, new exp.

## 2. Microphysics evaluation

Cloud liquid and ice (+snow and graupel) was evaluated against ACTRIS observations, using cloud-net algorithms (Hogan et al. 2004) against five ground-based remote sensors in central Europe for November, 2011.



Figure 4: Domain used for experiments, stations Chilbolton, Cabauw, Lindenberg, Leipzig and Juelich.

A comparison was made between Harmonie 38h.2 with AROME physics, and the option of switching off new cloud ice-microphysics (OCND2=FALSE, Ivarsson, 2014).

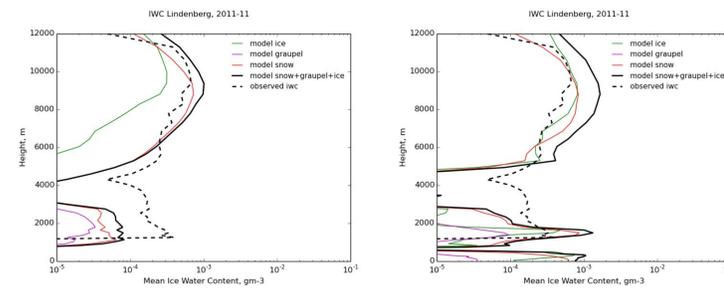


Figure 5: IWC for Lindenberg, November 2011. Left OCND2 = TRUE, Right OCND2 = FALSE.

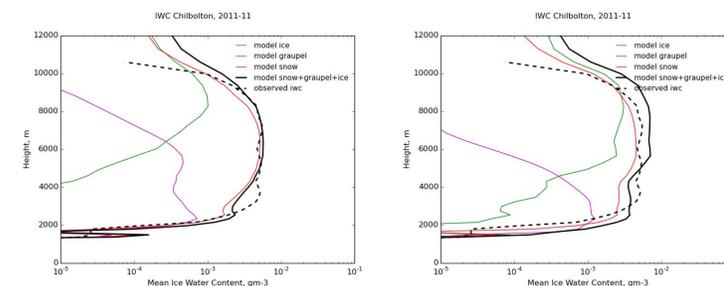


Figure 6: IWC for Chilbolton, November 2011. Left OCND2 = TRUE, Right OCND2 = FALSE.

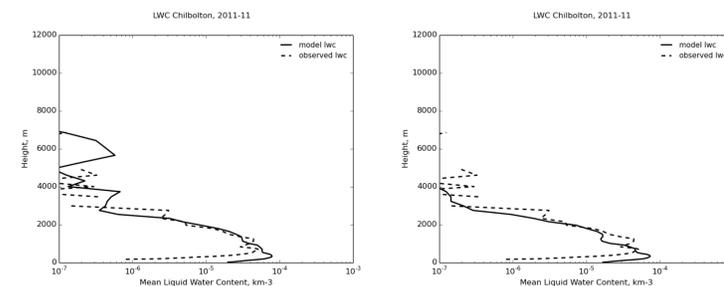


Figure 7: LWC for Chilbolton, November 2011. Left OCND2 = TRUE, Right OCND2 = FALSE.

For all stations, the new cloud microphysics option OCND2 gave a better comparison between the modelled and observed IWC for high clouds. In all cases, modelled cloud ice was reduced and modelled snow was increased with OCND2 = TRUE. For low level clouds there is too large of a reduction of ice+snow+graupel in case of OCND2 = TRUE.

The fit of liquid water is very good, although perhaps too much liquid around 6 km with option OCND2 = TRUE.

## 3. RACMO scheme in MetCoop

Results with the RACMO turbulence scheme in Harmonie-AROME-MetCoop are encouraging. In winter, the scheme gives an improvement in mean sea level pressure, lower troposphere temperature, and frequency bias of 12 h precipitation. In summer, an improved diurnal cycle of the 10 m wind is seen.

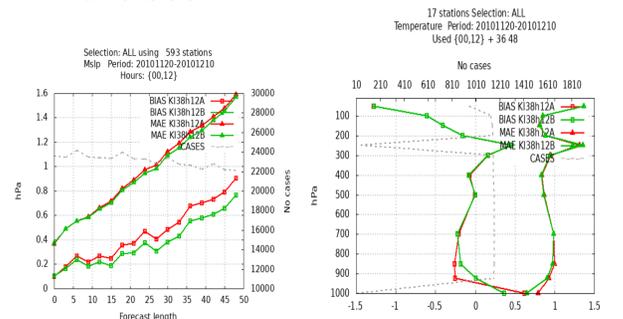


Figure 8 Left: Bias and MAE for mslp. Right: Bias and MAE temperature profiles. Reference (red), RACMO experiment (green).

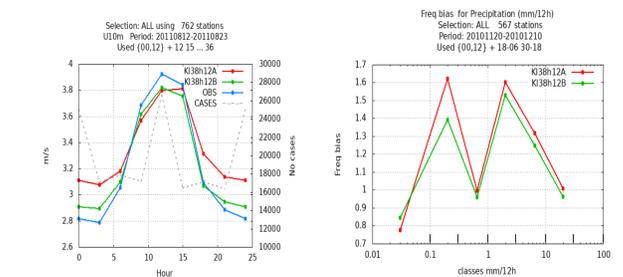


Figure 9: Left: diurnal cycle of 10 m wind, Right: frequency bias of 12 h acc precip. Red: reference, green RACMO exp and blue observations.

## 4. References

Ivarsson, 2014: Documentation of the OCND2 option in the ICE3 cloud- and stratiform condensation scheme in AROME.  
 Hogan, R. J., M. P. Mittermaier and A. J. Illingworth, 2004: The retrieval of ice water content from radar reflectivity factor and temperature and its use in the evaluation of a mesoscale model. *J. Appl. Meteorol.*  
 Kain, 2004: The Kain-Fritsch Convective Parameterization: An Update. *J. Appl. Meteor.*, 43, 170-181.