
Impact of the triggering function
on the diurnal cycle of deep convection
in SCM and GCM runs at ECMWF

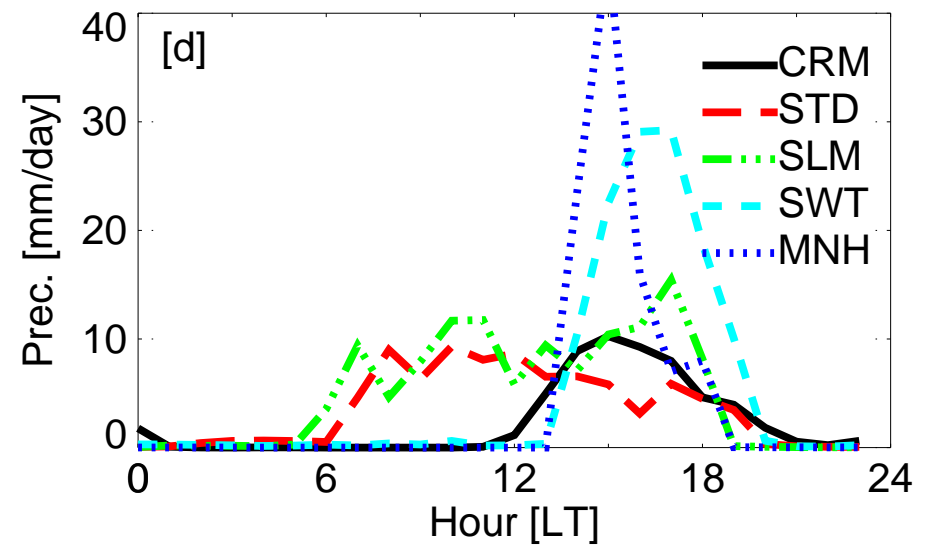
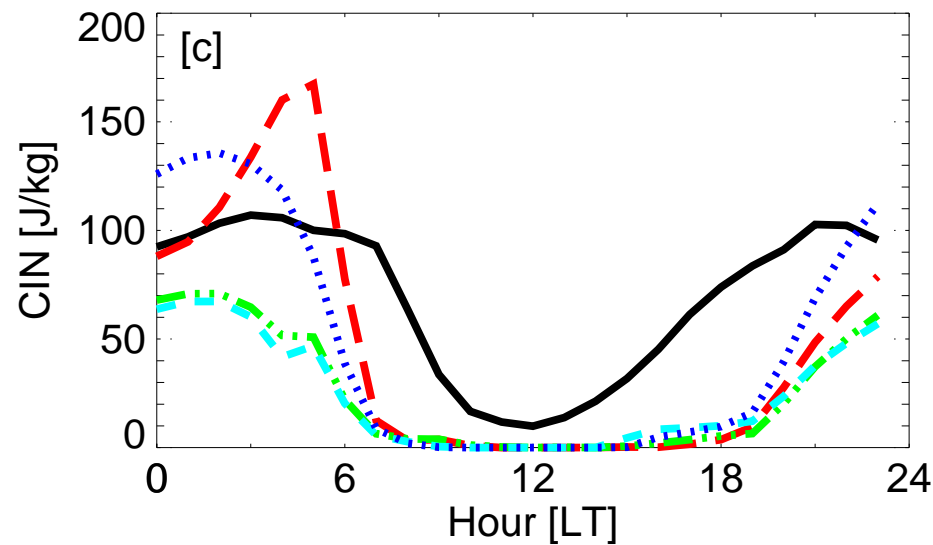
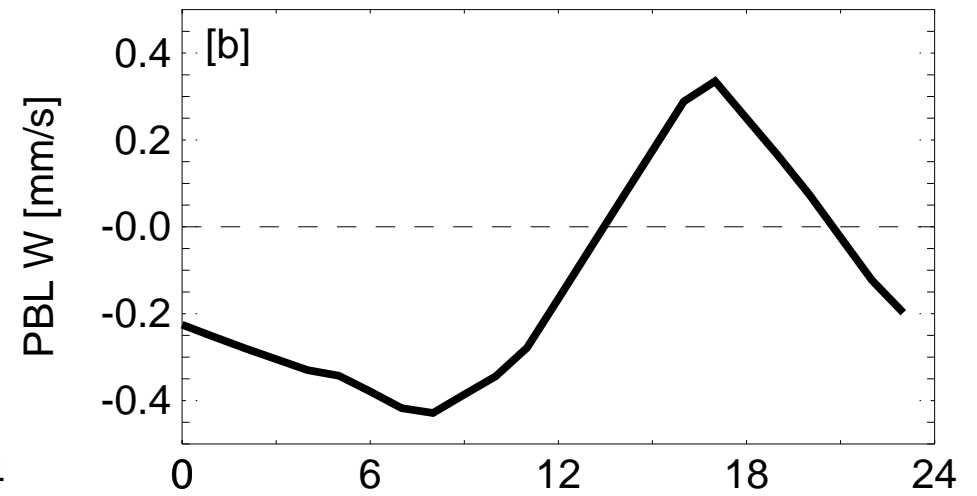
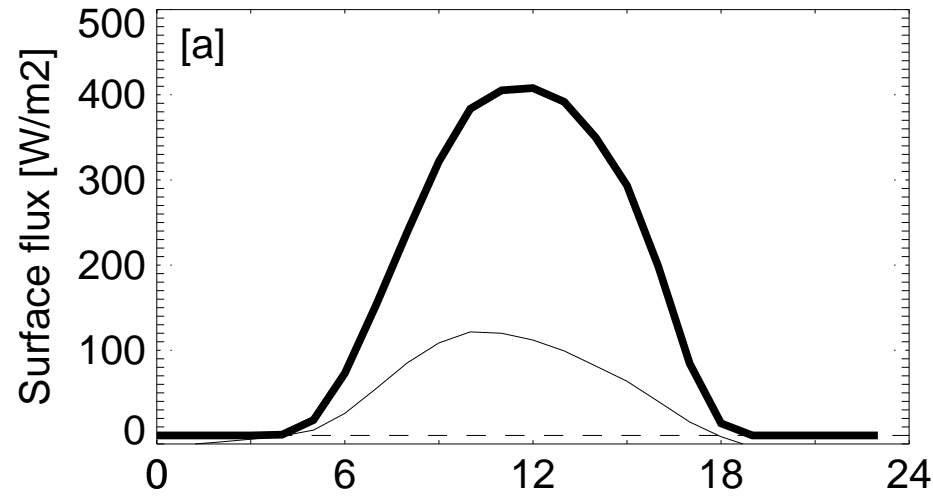
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¹CNRM-GAME, Toulouse and ²ECMWF, Reading

Short description of the convection schemes

Name	Trigger	Entrainment coefficient
STD	cloud depth > 200 hPa for surface parcel Tiedtke (1989), Jakob and Siebesma (2002)	$\epsilon = \epsilon_{\text{turb}} + \epsilon_{\text{org}}$ $\epsilon_{\text{turb}} = 4 \times 10^{-4}$ decreasing with height
SLM	cloud depth > 200 hPa for mixed-layer parcel in the lowest 350 hPa	idem STD
SWT	cloud depth > 200 hPa for mixed-layer parcel and positive buoyancy at cloud base	idem STD
MNH	cloud depth > 200 hPa for mixed-layer parcel and positive buoyancy at cloud base Bechtold et al. (2001)	$\epsilon_{\text{turb}} = 1.3 \times 10^{-4}$ constant with height

Idealized case study

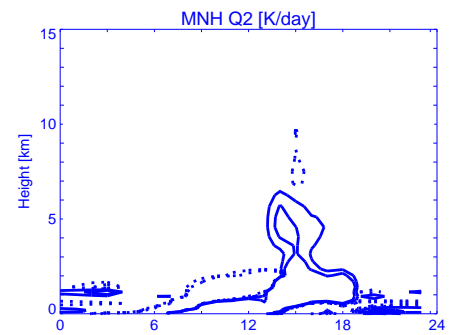
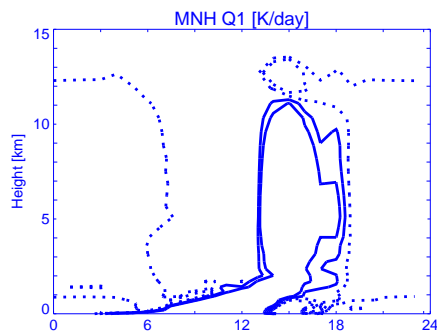
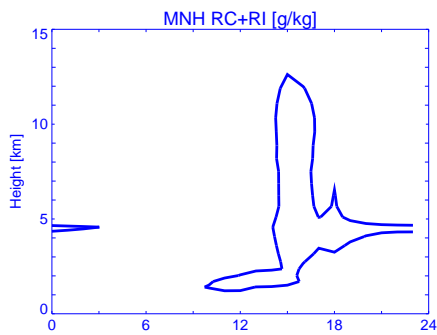
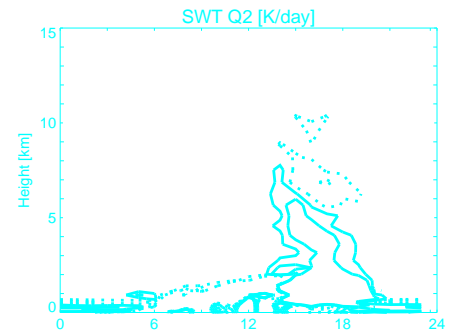
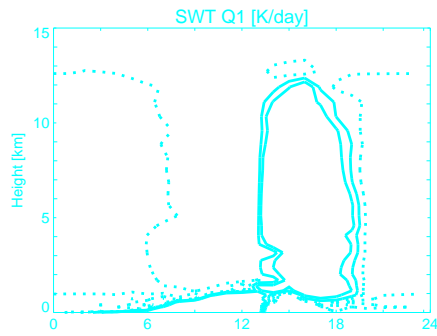
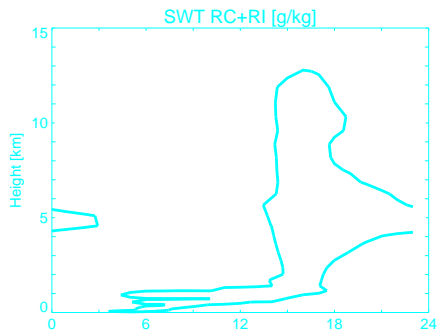
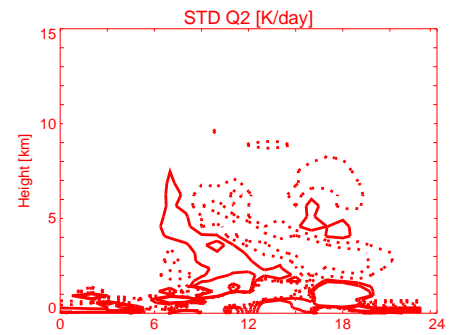
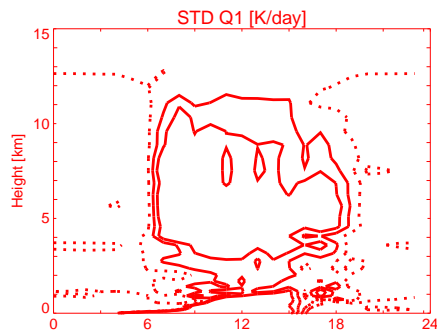
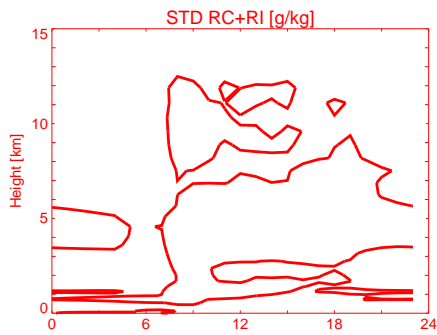
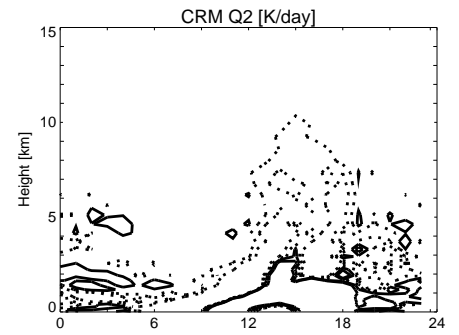
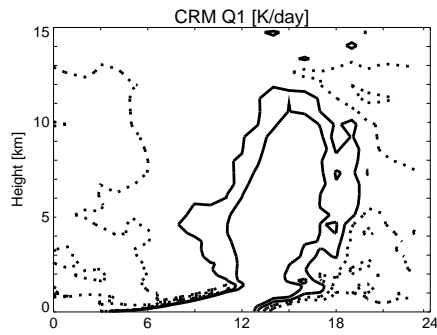
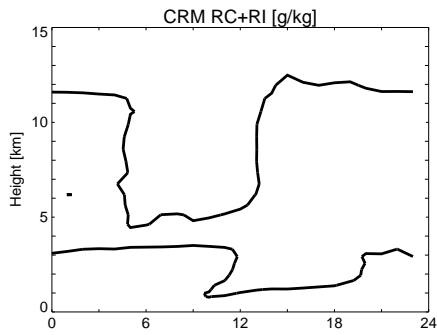


Idealized case study

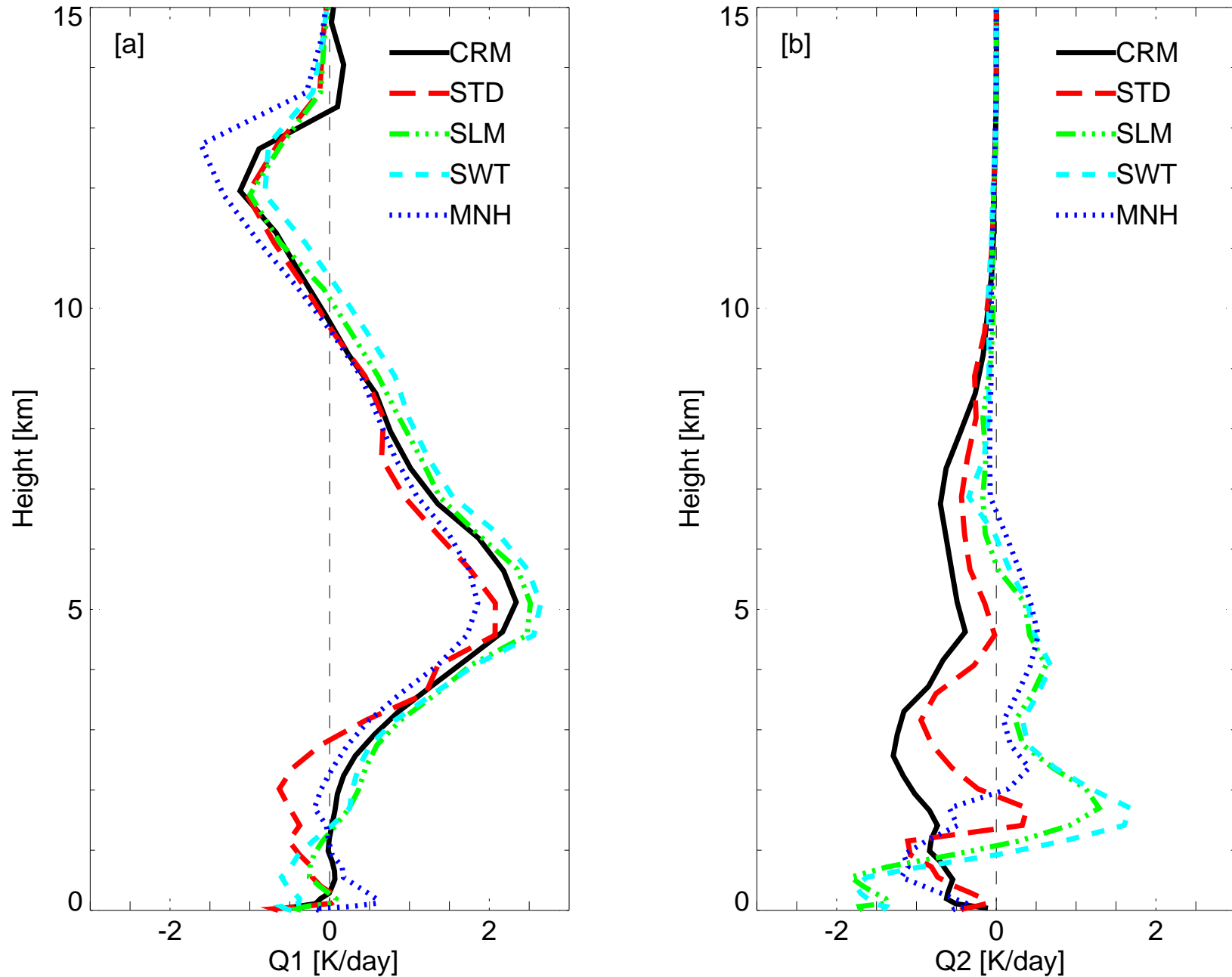
Cloud

Q1

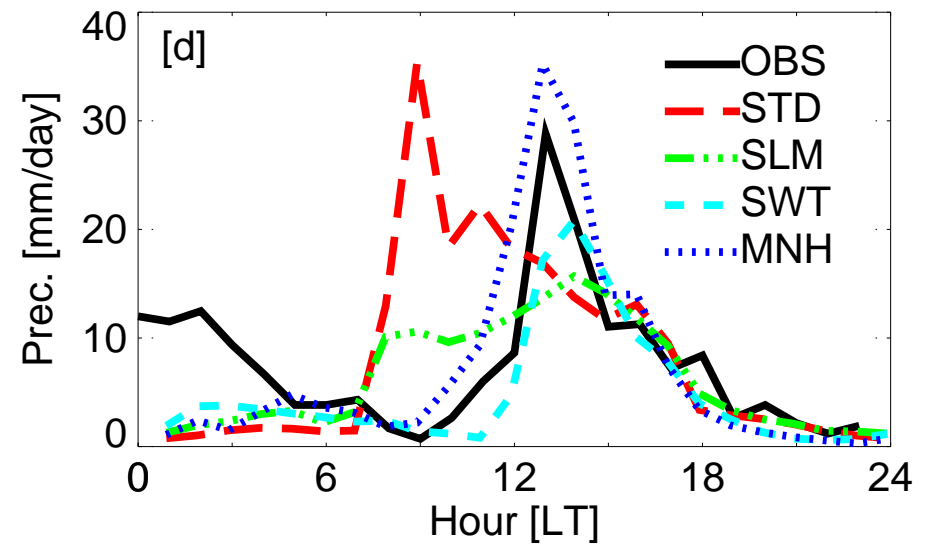
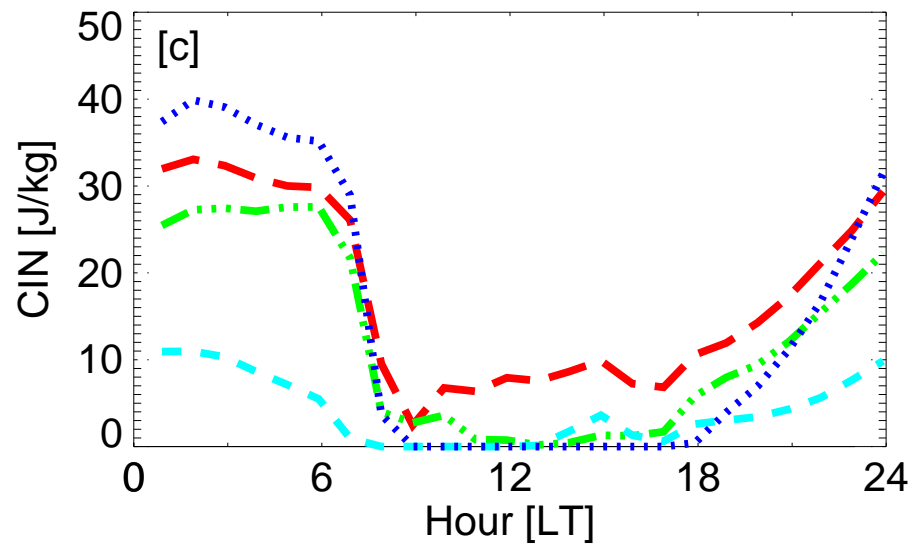
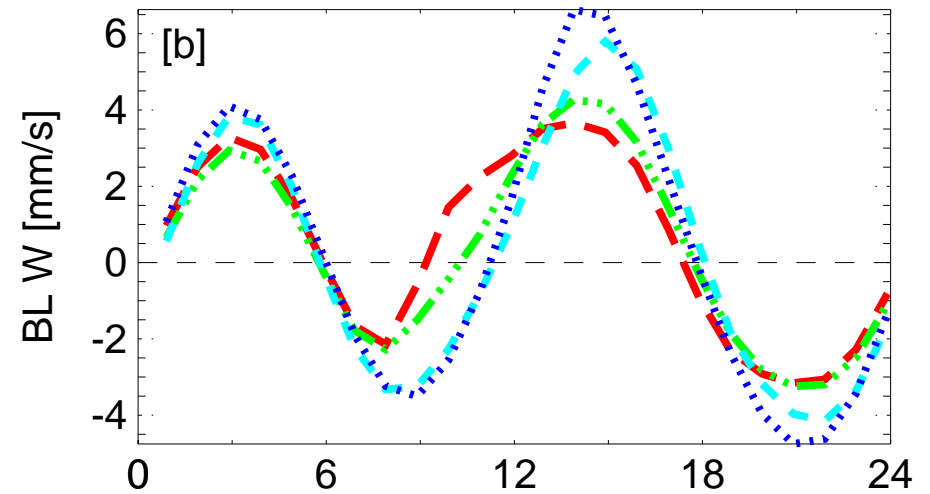
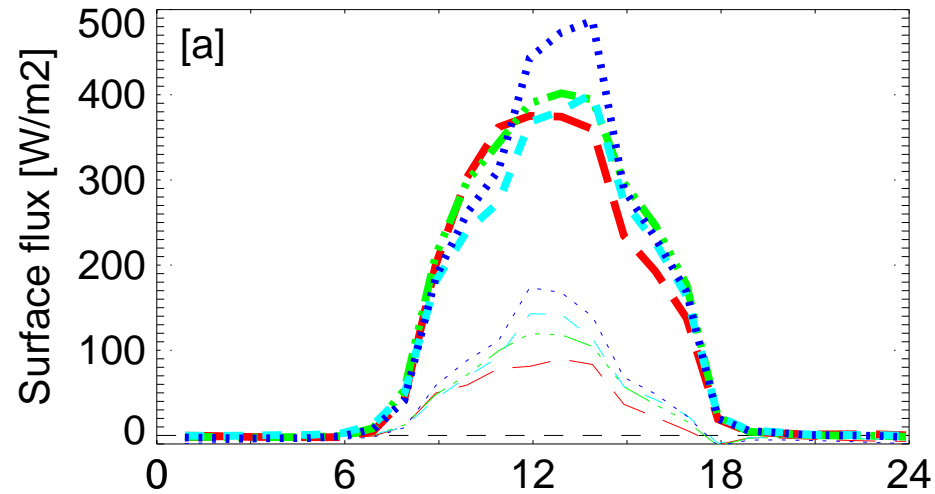
Q2



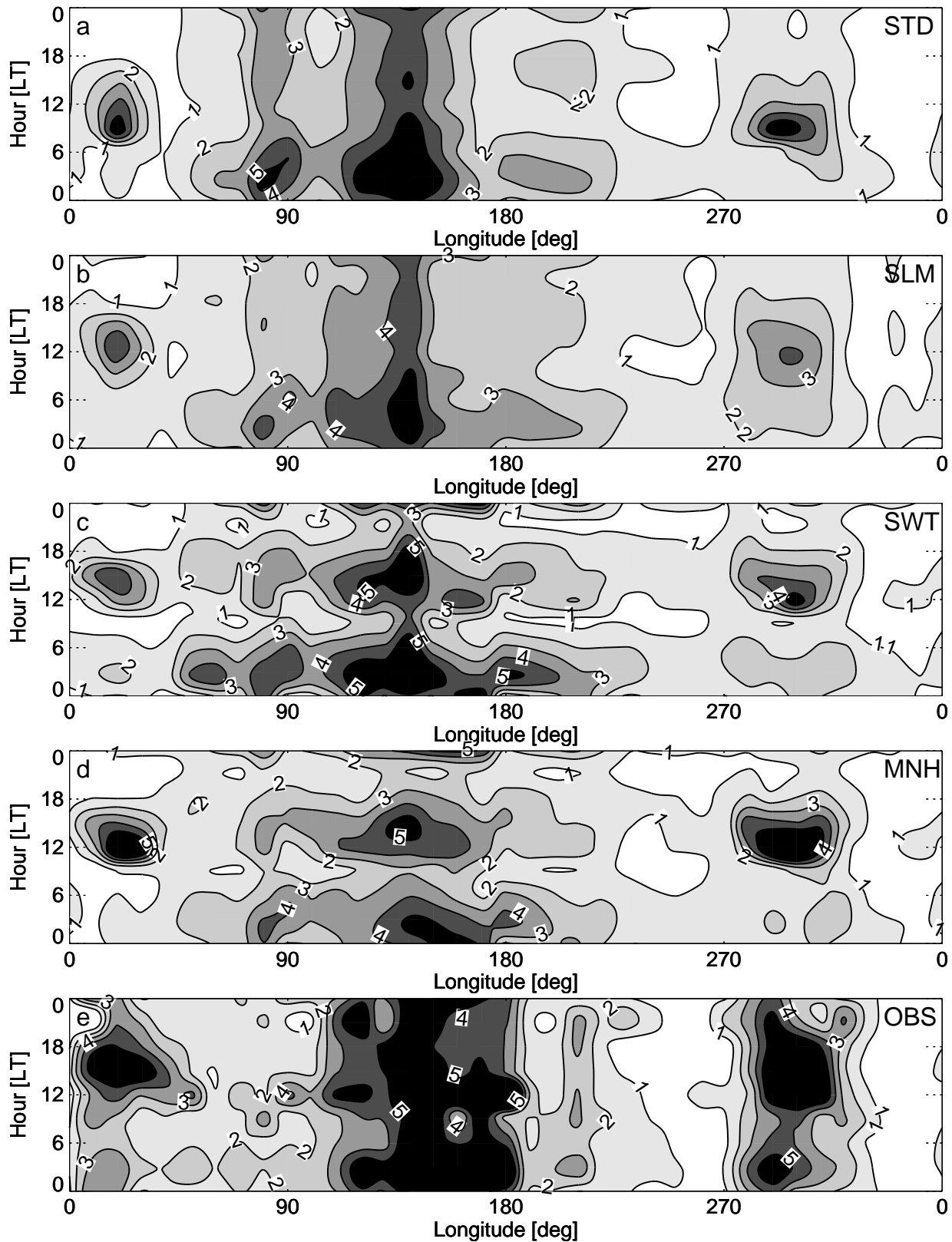
Idealized case study



Rondônia area - February 1999



Diurnal cycle at global scale



Precipitation (mm day⁻¹)

Fields averaged between 20°S and 20°N

Conclusion

**the current models can reasonably represent
the diurnal cycle of convective precipitation**

- ☞ when an appropriate trigger function is used in the convection scheme
- ☞ when the late morning shallow convection is correctly represented.

phase of tropical convection \leftrightarrow semidiurnal tide of large-scale convergence:
the trigger could be related to the large-scale convergence