Role of thermodynamic and turbulence processes on the fog life cycle during SOFOG3D experiment

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Context and objective

Research questions

- What processes contribute to the transition from stable to adiabatic fog?
- What processes contribute to the dissipation of fog?



Flights delayed



Pile-up on the Chaban bridge, Bordeaux

Tools:

- In-situ and remote sensing measurements during SOFOG3D field campaign
- Adiabatic fog conceptual model to derive additional key fog variables unobserved

•Data and studied area

Orography of the Study area



100 x 100 km around the supersite

5 weather stations around the supersite (Visibility, temperature, wind, etc.)

Visibility at the supersite, "La charbonnière"



Based on Tardif and Rasmussen (2007)

31 fog cases observed at Charbonnière, supersite (SS) during Nov 2019 – Mar 2020 period

- 4 heaviest fogs documented :
- (2) radiation
- · (2) radiation-advection

Visibility at the supersite



• Remote sensing at the supersite



Basta Mini Cloud, fog thickness



Windcube V2 Wind & TKE



Radiometer Hatpro LWP & stability



CL31 CBH – fog dissipation

Fog conceptual model

In-situ and remote sensing data T, P, visibility, LWP, CTH Fog adiabatic

conceptual model Toledo et al., 2021

 $\alpha_{eq} < 0$

z

СТН

Fog key parameters: equivalent adiabaticity & Reservoir

Equivalent adiabaticity by closure

$$\alpha_{eq}^{closure} = \frac{2 \left(LWP - LWC_0 CTH \right)}{\Gamma_{ad} (T, P) CTH^2}$$

Transition from stable to adiabatic fog

LWC is higher in the lower fog layers: Shallow stable fog



LWC increases with height \rightarrow Adiabatic fog \rightarrow Fog is transitioning from shallow to adiabatic **7**

Methodology Fog conceptual model



• Low fog LWP at SOFOG3D

Toledo et al., 2021

SIRTA

• More fog LWP at

 Equivalent adiabaticity by closure consistent at both sites – used to define the transition stable/adiabatic fog

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t1: start of the transition

- t2: end of the transition
- td: dissipation time

Fog formation, evolution and dissipation processes



- Formation: Cloud free East-West gradient of fog onset – low wind – cooling rate
- Transition: Increase in wind and temperature
- Dissipation: WS > 2 m/s

- Formation: Rain/stratus East-West gradient – low wind – cooling rate
- Transition: Increase in wind and temperature
- Dissipation: WS > 2 m/s

Fog formation, evolution and dissipation processes



- Formation: Easterly jet synoptic
- Stable : TKE < 0.2 m² s⁻² and σ_w^2 < 0.02 m² s⁻²
- Transition stable/adiabatic (TKE [0.2 $-0.4 \text{ m}^2 \text{ s}^{-2}$] and $\sigma_w^2 [0.02 0.04 \text{ m}^2 \text{ s}^{-2}]$, SHF [0 10 W m⁻²])
 - Dissipation by turbulence (TKE > $0.4 \text{ m}^2 \text{ s}^{-2} \text{ and } \sigma_w^2 \text{>} 0.04 \text{m}^2 \text{ s}^{-2} \text{ and SHF}$ > 10 W m⁻²]) – thermal and mechanic – RLWP (< 0) estimated an early (1h before) dissipation
- Formation: Westerly jet (Atlantic inflow) –
- Stable : TKE < 0.2 m² s⁻² and σ_{w}^{2} < 0.02 m² s⁻²
- Transition stable/adiabatic mechanical turbulence: (TKE $[0.2 - 0.4 \text{ m}^2 \text{ s}^{-2}]$ and $\sigma_w^2 [0.02 - 0.04 \text{ m}^2 \text{ s}^{-2}]$, SHF < 0 W m⁻²])
- Dissipation: southerly flow turbulence (TKE $\ge 0.3 \text{ m}^2 \text{s}^{-2}$ and $\sigma_w^2 > 0.04 \text{ m}^2 \text{s}^{-2}$) – thermal and mechanic – RLWP (< 0) estimated the dissipation time

Fog formation, evolution and dissipation processes



- Transition stable/adiabatic more marked in radiation fog – lowering of the inversion top height
- Link between temperature inversion strength and fog lifetime

Summary

- Radiation fog cases have longest lifetime (more than 12 hours) linked to very cold atmospheric conditions associated with a continental easterly nocturnal low-level jet stable/adiabatic fog transition driven by advection (TKE [0.2 0.4 m² s⁻²] and σ_w^2 [0.02 0.04 m² s⁻²], SHF [0 10 W m⁻²]) dissipation in daytime by thermal and mechanical turbulence (TKE > 0.4 m² s⁻² and σ_w^2 > 0.04 m² s⁻² and SHF > 10 W m⁻²)
- Advection-radiation case studies have shortest lifetime linked to the low surface boundary layer stability due to the vertical mixing generated by the westerly strong wind –transition phase is driven by advection (TKE $[0.2 0.4 \text{ m}^2 \text{ s}^{-2}]$ and $\sigma_w^2 [0.02 0.04 \text{ m}^2 \text{ s}^{-2}]$, SHF < 0 W m⁻²) the dissipation phase is driven by night-time warm air advection generating mechanical turbulence (TKE at least 0.3 m² s⁻² and $\sigma_w^2 > 0.04 \text{ m}^2 \text{ s}^{-2}$).
- This study also demonstrates the importance of using instrumental synergy (with microwave radiometer, wind lidar, weather station, and cloud radar) and a fog conceptual model to better predict fog characteristics and dissipation time at nowcasting ranges.



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